Full Research Article

Distribution of Different Forms of Soil Acidity in Selective Laterite Soils under Cashew **Plantations in Coastal Karnataka**

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Abstract

A study was conducted to examine soil morphological, physico-chemical properties and different forms of soil acidity of six selected laterite soils under cashew plantations in coastal Karnataka. Pedons were dug to a depth of 2 m at six selected locations viz. Pala hillock, DCR-Shanthigodu, Muchipadavu, Mulia, ARS-Ullal and Palikudel and horizon wise samples were collected to study the various soil parameters including different forms of soil acidity. Cation exchange capacity and base saturation of the soils were varied from 7.60 to 19.8 cmol (p⁺) kg⁻¹ and 4 to 32%, respectively. The exchangeable acidity varied from 0.06 to 2.20 cmol (p+) kg-1 and extractable acidity ranged from 8.16 to 47.89 cmol (p+) kg⁻¹. Presence of low exchangeable acidity could be due to efficient recycling of basic cations and higher exchangeable acidity in few locations may be due to the presence of exchangeable ions like H⁺ and Al³⁺. Results obtained from this study indicated that the Pedons 1, 4, 5 and 6 requires more lime than others to reclaim for maximizing cashew production by restoring the available soil nutrients.

1. Introduction

Cashew is a major export crop in India with highest percentage of exports coming from Western parts of coastal India, majorly from coastal Karnataka. However, the production has dwindled in recent decades due to variable climatic and soil constraints (Bhat, 2007). From the coastal regions cashew was introduced to the plains of Karnataka, Chhattisgarh and in the hilly areas of Eastern states. Although area under cashew has been steadily increasing in India in the recent year, its production is not satisfactory in most part of the country. Cashew is mostly grown on laterite and red soils and coastal sandy in the states, which are very poor in fertility and the yields are low in these soils unless the trees are fertilized regularly (Anon, 2003). Major problems are shallow depth, coarse texture and severe acidity. Soil acidification and related land degradation are becoming more important worldwide and are challenging the concept of sustainability of current land management system (Noble et al., 2000).

Acid soils cover about 49 mha, approximately 34% of cropped land in India (Maji et al., 2012). The acid soils (pH <6.5) occupy approximately 90 mha of the geographical area (Sharma and Sarkar, 2005). Strongly acidic soils (below pH 5.5) occupy

around 25 mha of arable land in India. Such type of problematic soils in India are mainly concentrated in North-Eastern (NE) region and Western Ghats, Jharkhand, Chhatisgarh, Uttarakhand with sporadic distribution in Himachal Pradesh, Orissa and West Bengal, etc (Panda, 1987). Soil acidity has been recognized as an important agricultural problem, which adversely affects the crop production, directly or indirectly. Cashew production involves a complex interaction between the environment, soil parameters, and nutrient dynamics. Because of this fact, the soil must be studied in terms of the productive potentials. Failure to understand these complexities has resulted in lack of good crop production and management techniques; hence agricultural production has tended to be low. Soil fertility decline is considered as an important cause for low productivity of many soils (Sanchez, 2002). Assessing soil fertility decline is difficult because most soil chemical properties either change very slowly or have large seasonal fluctuations. This decline includes; nutrient depletion, nutrient mining, acidification (decline in pH and or an increase in exchangeable Al), loss of organic matter and increase in toxic elements (e.g. Al, Mn) (Hartemink, 2006). Soil acidity constitutes one of the important factors, which influence the chemistry as well as fertility management of tropical soils like laterite soils. These acidity

parameters will help us in understanding the rate at which the process of laterisation proceeds in a particular land use system (Chenithung et al., 2014). Since, not much study has done in the laterite soils of coastal Karnataka on different forms of soil acidity effect on cashew production. Hence, the study was aimed at understanding the various soil acidity parameters of six selected laterite soils along with other soil properties under cashew from coastal Karnataka.

2. Materials and Methods

2.1. Study area selection

A case study was undertaken in four taluks of Dakshina Kannada district of Coastal Karnataka in during 2012–13 (Table 1). The geology of the area consists of rock formation of archean complex (archean granites and gneisses) with sparse bands of dharwar system of rocks which are capped by laterite at many places. The study area is included in the coastal zone of Karnataka that comprises two major geomorphological zones viz. Fluvio-littoral formation and dissected hilly hinter land (Dasog and Patil, 2002).

2.2. Climatic conditions of the study area

Study area represents hot and humid climate with annual precipitation ranged from 3592 to 3842 mm. The maximum and minimum temperature of the study area recorded was 36 °C and 20 °C respectively. Relative humidity varied from 76 to 97%, due to low altitude. Mean annual potential evaporation was also very high (>1000) especially in the coastal region. The rainfall distribution during May to November month was frequent, with heavy rainfall recorded during July in all the

Table 1: Major	cashew gr	owing areas in the coastal	Karnataka
Taluk	Village	Land use	Yield
			(kg ha ⁻¹)
1. Puttur	Pala hillock	Forest with plantation of cashew	725
2. Puttur	DCR- Shanth- igodu	Well maintained cashew farm with intercrops of pineapple and onion	690
3. Bantwal	Muchi- padavu	Poorly maintained cashew plantation with other tree species	685
4. Bantwal	Mulia	Degraded cashew plantation in the hillock	461
5. Manga- lore	ARS- Ullal	Well maintained cashew farm	710
6. Karkala	Paliku- del	Patches of cashew plantation with tree species and shrubs	425

study location. Length of dry season is less about four months and length of crop growing period is about 190 to 210 days. Water balance diagrams were prepared using method suggested by Thornthwaite and Mather (1955) and model of Higgins and Kassam (1981), which is used for estimating of water deficit, water surplus and moisture index, which could be useful to determine the water surplus and deficient periods (Figure 1).

2.3. Soil sampling and analysis

Soil samples collected from horizon wise of all profiles were dried in air and then crushed in wooden mortar with a wooden pestle and sieved through 2 mm sieve to separate the coarse fragments (>2 mm). The fine earth samples were stored in separate containers and used for analysis. Particlesize analysis of the samples was carried out by international pipette method. Electrical conductivity, pH, cation exchange capacity and base saturation were determined using standard methods (Jackson, 1973). Exchangeable Hydrogen and Aluminium together constitute the exchangeable acidity of the soil. These exchangeable acid cations were extracted using 1.0 N KCl and total exchangeable acidity was determined by titrating with standard alkali (NaOH) solution. Exchangeable aluminium was determined by treating the titrated solution with excess potassium fluoride and then titrating with standard sulphuric acid until the pink colour disappears. Exchangeable hydrogen was calculated by difference (Sarma et al., 1987). Extractable acidity was determined by using barium chloridetriethanolamine. The soil buffer mixture was shaken in an end to end shaker for 16 hours, the contents centrifuged and the clear supernatant was titrated against standard HCl using bromocresol green as indicator (Sarma et al., 1987). Lime requirement was done by treating soil with a SMP buffer (1.8 g paranitrophenol, 2.5 ml of TEA, 3 g of potassium chromate, 2 g calcium acetate and 53.1 g of CaCl₂ dissolved in one litre of water with pH 7.5) solution and the lime requirement was calculated from the reduction in pH of the buffer (Shoemaker et al., 1961).

3. Results and Discussion

3.1. Soil morphological characteristics

Brief morphological features of the pedons are presented in Table 2. The solum depth varied from deep to very deep. Pedons had their Munsell colour notation in the hue of 10YR/7.5YR/5YR/2.5YR with value 3 to 4 and chroma 4 to 8. The soil colour appears to be the function of chemical and mineralogical composition as well as textural make up of soils and conditioned by topographic position and moisture regime (Shiva prasad et al., 1998). Pedons 1, 2, 4 and 6 had dark brown colour in surface soils and dark red to red colour in sub soils. In pedon 3 and 5, surface soil colors were yellowish red and dark reddish brown and whereas the sub soil colour was

red. All the layers of various pedons contained high amount of gravel percentage, which increased with increasing depth. In all the pedons texture varied from sandy clay loam to clay and clay content were varied from 24.5 to 66.4%. Low clay

content in the surface horizons might be due to the illuvation or translocation of clay from the surface to subsurface horizons. The wide textural variation might be due to variation in soil formation, in situ weathering. Clay skin or cutans distribution

Depth	Horizon	Sand	Silt	Clay	Texture	Moist Co-	Coarse	BD	PD	Pores	
(cm)		(2.0- 0.05)	(0.05- 0.002)	(<0.002)	class (USDA)	lour (Matrix)	frag- ments	Mg m ⁻³		paces (%)	
Pedon 1 (Loamy- ske			hermic Ustic							
0-21	A1	58.9	15.6	25.5	scl	7.5YR 3/4	10	1.42	2.48	43	
21-41	Bt1	55.1	9.4	35.5	sc	7.5YR 3/4	20	1.33	2.57	48	
41-68	Bt2	62.2	6.4	31.4	scl	5YR 3/4	50	1.38	2.60	47	
68-95	Bt3	57.8	9.9	32.3	scl	2.5YR 3/6	40	1.38	2.64	48	
95-123	Bt4	59.5	11.3	29.2	scl	2.5YR 3/6	60	1.57	2.69	42	
Pedon 2 (Fine, mixed	, isohyper	thermic Ox	ic Dystruste	ots)						
0-19	Ap	38.6	8.9	52.5	c	7.5YR 3/4	30	1.44	2.54	43	
19-42	Bw1	36.8	12.6	50.6	c	2.5YR 3/6	30	1.49	2.52	41	
42-71	Bw2	33.6	15.3	51.2	c	2.5YR 3/6	80	1.45	2.52	42	
71-90	Bc	44.0	13.9	42.1	c	2.5YR4/6	80	1.50	2.59	42	
Pedon 3 (Fine, mixed	, isohyper	thermic Ox	ic Dystruster	ots)						
0-25	Ap	26.7	10.2	63.0	c	5YR 4/6	35	1.27	2.56	50	
25-50	Bw1	21.4	13.6	65.1	c	2.5YR 3/6	25	1.21	2.55	52	
50-84	Bw2	23.3	14.6	62.1	c	2.5YR 3/6	20	1.13	2.49	54	
84-128	Bw3	16.3	17.3	66.4	c	2.5YR 3/6	15	1.13	2.48	54	
128-160	Bw4	21.2	20.5	58.3	c	2.5YR 3/6	15	1.14	2.49	54	
Pedon 4 (Clayey-skel	etal, mixe	d isohypert	hermic Ustic	Haplohum	ults)					
0-17	A	37.2	22.2	40.6	c	7.5YR 3/4	40	1.31	2.48	47	
17-41	Bt1	28.0	19.6	52.4	c	7.5YR 3/4	60	1.23	2.48	50	
41-60	Bt2	29.4	9.4	61.2	c	2.5YR3/6	70	1.36	2.50	45	
60-90	Bt3	37.1	5.0	57.9	c	2.5YR 4/6	70	1.23	2.54	51	
90-155	BC	50.3	10.9	38.8	c	2.5YR4/6	70	1.20	2.51	52	
Pedon 5 (Clayey-skel	etal, mixe	d isohypert	hermic Typic	c Rhodustu	lts)					
0-20	Ap	61.0	6.8	32.2	cl	5YR 3/4	20	1.33	2.58	48	
20-40	Bw	56.8	7.4	35.8	sc	2.5YR 3/6	40	1.36	2.62	48	
40-70	Bt1	49.6	6.2	44.2	sc	2.5YR 3/6	60	1.52	2.52	40	
70-99	Bt2	41.1	6.8	52.1	c	2.5YR 3/4	60	1.60	2.53	37	
99-127	Bt3	46.8	5.0	48.2	sc	2.5YR 4/8	70	1.60	2.49	36	
127-165	Bt4	50.3	4.3	45.4	sc	2.5YR 4/8	60	1.58	2.53	37	
Pedon 6 (Clayey-skel	etal, mixe	d, isohyper	thermic Usti	c Palehumu	ılts)					
0-28	A	66.8	8.7	24.5	scl	7.5YR 3/4	30	1.30	2.50	48	
28-55	Bt1	50.5	12.3	37.2	sc	5YR 3/6	35	1.44	2.60	45	
55-90	Bt2	48.7	11.2	40.1	sc	2.5YR 3/6	40	1.44	2.65	45	
90-130	Bt3	43.5	12.5	44.0	c	2.5YR 4/6	40	1.47	2.57	43	
130-180	Bt4	43.5	11.3	45.2	c	2.5YR 4/6	50	1.43	2.68	46	

of the profiles were unable to notice due to high rainfall in the study area. Based on the laboratory analysis on clay content, CEC and base saturation, it was found that the pedons 1, 4, 5 and 6 had argillic endopedons and pedon 2 and 3 had cambic endopedons. The bulk density, particle density and porosity ranged from 1.13 to 1.67, 1.30 to 1.69 Mg m⁻³ and 36 to 54 % respectively. The irregular trend of porosity with depth was due to the illuviation and eluviation processes of finer fractions in to different horizons (Chadhaand Nair, 1998; Bhattacharyya et al., 2009).

3.2. Physico-chemical characteristics

The detailed physico-chemical characteristics of the soils are given in Table 3. All the pedons were moderately to strongly acidic with pH ranging from 4.8 to 6.0. The acidic nature of

the soil might be attributed mainly to the leaching of the bases due to the existing high rainfall conditions and to some extend may contributed from acidic parent materials. (Shivaprasad et al., 1998) while characterizing the soils of Karnataka observed that the soils derived from granite-gnesis parent materials were found to be slightly acidic to near neutral in soil reaction. The organic carbon content of the soils was found to be high in surface soils and low in sub surface soils. The organic carbon content was decreased with increasing depth. This is attributed to the addition of plant residues and farmyard manure to surface horizons (Sivasankaran et al., 1993). Cation exchange capacity and base saturation varied from 7.6 to 33.6 cmol (p+) kg-1 and 7 to 34% respectively, which showed decreasing trend with increasing depth of soil (Dey and Sehgal, 1997; Patton et al., 2007).

Depth	Ho-	sico-chemical and acidity characteristics of the soils pH (1:2.5) EC OC CEC BS Ex- Exchangeable Requ												quirement			
(cm)	ri-	1	p11 (1.2.3)		(dS/m)	(%)	(cmol (p+) kg ⁻¹)			(%)	tract-	acidity			of lime to		
	zon					()	(1, 6, 7			(70)	able acidity (BaCl ₂ -	1.0 M KC ₁			bring the soil to desired pH (t ac ⁻¹)		
		H_2O	1.0M				NH ₄ OAc	Sum	ECEC		TEA	$H^{\scriptscriptstyle +}$	$A1^{3+}$	То-	6.0	6.4	6.8
			KCl	CaCl ₂			pH 7.0	of cation			pH 8.2)			tal			
							7.0			_	cmol (p+) kg ⁻¹						
							Ustic Hap										
0-21	A1	5.4	4.2	4.5	0.05	3.07	15.5	51.79	4.23	25	47.8	0.55		0.88		5.2	6.0
21-41	Bt1	5.3	4.0	4.1	0.01	1.53	14.7	33.04	2.54	7	31.9	0.45	1.46	1.91	3.9	4.7	5.4
41-68	Bt2	5.6	4.0	4.2	0.01	0.38	13.1	19.59	2.30	12	17.9	0.39	0.61	1.00		2.2	
68-95	Bt3	5.6	4.1	4.3	0.01	0.32	11.9	18.14	1.49	10	16.9	0.25	0.27	0.52	1.0	1.2	1.4
95-123	Bt4	5.7	4.4	4.6	0.01	0.13	7.7	13.04	1.46	16	11.7	0.0	0.0	0.0	0.0	0.0	0.0
Pedon-2				-													
0-19	Ap	5.5	4.3	4.4	0.02	1.68	12.4	28.01	3.88	31	24.4	0.0	0.0	0.0	2.3		3.1
19-42	Bw1	5.6	4.5	4.6	0.01	0.97	11.0	21.03	1.43	13	19.6	0.0	0.0	0.0	1.0	1.2	1.4
42-71	Bw2	5.7	4.0	5.0	0.01	0.70	8.6	17.60	1.64	19	15.9	0.0	0.0	0.0	0.0	0.0	
71-90	Bc	6.0	5.3	5.3	0.01	0.14	7.6	16.43	2.0	26	14.4	0.0	0.0	0.0	0.0	0.0	0.0
Pedon-3				•													
0-25	Ap	5.5	4.2	4.2	0.01	0.97	10.2	23.72	2.59	20	21.63	0.33	0.50			1.7	
25-50	Bw1	5.3	4.3	4.3	0.01	0.82	8.0	26.44	3.04	34	23.73	0.28	0.33	0.61	1.0	1.2	1.4
50-84	Bw2	5.5	4.5	4.4	0.02	0.50	7.8	22.29	1.41	18	20.88	0.0	0.0	0.0	1.0		1.4
84-128	Bw3	5.3	4.5	4.3	0.00	0.40	10.2	20.19	1.37	13	18.82	0.0	0.0	0.0			
128-160	Bw4	5.4	4.7	4.4	0.03	0.26	9.8	15.99	1.39	14	14.60	0.0	0.0	0.0	0.0	0.0	0.0
							Ustic Hap				42.0						
0-17	A	5.5	4.3	4.4	0.29	2.80	19.8	44.4	3.5	11	42.0	0.21	1.23	1.44		3.7	
17-41	Bt1	5.4	4.2	4.3	0.05	2.04	14.8	40.8	3.2	15	38.6	0.77	0.96	1.73	3.5	4.2	
41-60	Bt2	5.5	4.2	4.3	0.01	1.42	14.3	31.8	3.6	18	29.3	0.12	1.07	1.19	3.1	3.7	
60-90	Bt3	5.5	4.3	4.4	0.02	0.65	9.1	30.7	3.3	32	27.8	0.42	0.39	0.81	1.4		1.9
90-155	BC	5.3	4.5	4.6	0.04	0.43	10.7	21.8	1.9	18	19.8	0.0	0.0	0.0		1.7	

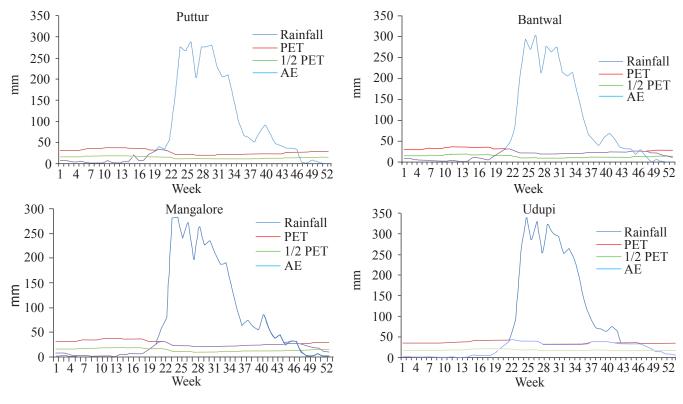
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Depth	Но-	pH (1:2.5)		EC	OC		CEC		BS	Ex-	Exchangeable			Requireme			
(cm)	ri-				(dS/m)	(%)	(cmol	(p+) kg	g-1)	(%)	tract-		acidity		of	lime	to
	zon										able	$1.0 \mathrm{M}\mathrm{KC}_{1}$			bring the soil		
											acidity				to d	esirec	l pH
					_					_	(BaCl ₂ -				((t ac-1)
		H ₂ O	1.0M	0.01M			NH ₄ OAc	Sum	ECEC		TEA	H^{+}	$Al^{\scriptscriptstyle 3+}$	To-	6.0	6.4	6.8
			KCl	CaCl ₂			pН	of			pH 8.2)			tal			
							7.0	cation			cmol						
Pedon-5	(Claye	ey-ske	letal, n	nixed isc	hyperth	ermic	Typic Rho	dustult	s)								
0-20	Ap	4.9	4.0	4.1	0.20	1.56	11.3	23.9	2.6	9	23.0	0.57	1.61	2.18	3.9	4.7	5.4
20-40	Bw	4.8	4.0	4.1	0.08	0.83	10.3	19.9	1.8	4	19.4	0.62	1.39	2.01	3.5	4.2	4.8
40-70	Bt1	5.1	3.9	4.1	0.25	0.13	12.0	20.5	3.1	12	19.0	0.58	1.62	2.2	3.1	3.7	4.2
70-99	Bt2	5.2	4.0	4.3	0.58	0.45	13.8	20.6	3.7	21	17.6	0.52	0.73	1.25	2.7	3.2	3.7
99-127	Bt3	5.3	4.1	4.3	0.27	0.36	11.2	16.8	3.5	25	13.9	0.39	0.61	1.00	1.8	2.2	2.5
127-165	Bt4	5.1	4.1	4.4	0.04	0.27	10.4	26.7	2.5	21	24.5	0.14	0.39	0.53	1.8	2.2	2.5
Pedon-6	(Claye	ey-ske	letal, m	nixed, is	ohyperth	ermic	Ustic Pale	humult	s)								
0-28	A	5.7	4.5	4.5	0.62	1.79	11.8	18.6	3.3	28	15.2	0.0	0.0	0.0	1.8	2.2	2.5
28-55	Bt1	5.5	4.4	4.4	0.35	1.30	11.2	18.0	2.4	20	15.7	0.26	0.14	0.40	2.3	2.7	3.1
55-90	Bt2	5.6	4.4	4.5	0.04	0.73	8.8	12.9	2.2	25	10.6	0.0	0.0	0.0	2.3	2.7	3.1
90-130	Bt3	5.4	4.8	4.8	0.08	0.36	9.4	10.7	2.5	27	8.1	0.0	0.0	0.0	1.0	1.2	1.4
130-180	Bt4	5.6	4.9	4.8	0.56	0.28	9.5	13.1	2.4	25	10.7	0.0	0.0	0.0	0.0	0.0	0.0

3.3. Soil acidity characteristics

Exchangeable acidity and extractable acidity varied from 0.06 to 2.20 cmol (p+) kg^{-1} and 8.16 to 47.89 cmol (p+) kg^{-1} .

The values were decreased with increasing depth of the all the soil profiles, more acid extract was found in less than 40 cm depth of soil (Table 3). Low exchangeable acidity could



be due to efficient recycling of basic cations which facilitates the soil to maintain a higher pH and low exchangeable acidity and aluminium (Talashilkar et al., 2006; Amenla et al., 2010). Higher exchangeable acidity in few locations could be due to the presence of exchangeable ions like H⁺ and A¹³⁺ which might have contributed towards soil acidity (Laxminarayana, 2010). The most important soil factors which control the different kinds of soil acidity are pH, organic matter content and exchangeable and extractable aluminium (Sharma et al., 1990). Rainfall is the most important factor contributing towards removal of bases. Higher rainfall and free drainage favoured removal of bases to a maximum level and its deposition in the lower horizons contributed to the build up of significant amount of calcium in the subsoils.

Soil pH was measured potentiometrically in H₂O and 1 mol L-1 KCl suspensions (1:2.5 soil: solution ratio), and the difference $\Delta pH = pH KCl - pH H_2O$ was used to estimate the sign of the net charge (Mekaru and Uehara, 1972; Steiner and Lana, 2013). The ΔpH indicated that these soils contained appreciable quantity of clay minerals with relatively constant surface charges (Bleekar and Sageman, 1990). According to Van Wambede (1992), the amount of extractable Fe is usually greater in Ultisols than Oxisols. In Ultisols, the major source of exchangeable and titratable acidity were monomeric Al3+ and Al hydroxyl compounds (Pal et al., 2014). The less extractable acidity indicated the less degree of weathering and leaching of the profile (Natarajan, 1995).

3.4. Lime management

Based on the distribution of soil acidity, lime requirement was calculated and is given in Table 3. Results revealed that pedons 1, 4, 5 and 6 requires more lime to neutralize surface soils than subsurface, whereas pedons 2 and 3 needs less lime. (Badrinath et al., 1997) pointed out that the current levels of cashew nut yield per tree are low and might be due to several factors which are limiting the economic yield. Hence, the application of inorganic fertilizers carrying P, K and Boron after the soil amelioration with dolomite lime stone will not only raises the soil pH but also favours the supplied plant nutrients in the available form to the plants (Laxminarayana, 2010).

4. Conclusion

The nature of soil acidity is mostly clay and pH-dependent. Electro statistically bound Al³⁺ is the major contributing factor towards the exchange and extractable acidity. Exchange acidity showed highly significant relation with exchangeable Al³⁺ rather than with exchangeable H⁺. Thus could be managed through proper cultural practices, addition of lime materials and base fertilizers, which will restore nutrients availability and sustain the cashew production in coastal Karnataka.

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